



Reviving the “Ganges Water Machine”: where and how much?

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Abstract. Runoff generated in the monsoon months in the upstream parts of the Ganges River basin (GRB) contributes substantially to downstream floods, while water shortages in the dry months affect agricultural production in the basin. This paper examines the potential for subsurface storage (SSS) in the Ganges basin to mitigate floods in the downstream areas and increase the availability of water during drier months. The Soil and Water Assessment Tool (SWAT) is used to estimate “sub-basin” water availability. The water availability estimated is then compared with the sub-basin-wise unmet water demand for agriculture. Hydrological analysis reveals that some of the unmet water demand in the sub-basin can be met provided it is possible to capture the runoff in sub-surface storage during the monsoon season (June to September). Some of the groundwater recharge is returned to the stream as baseflow and has the potential to increase dry season river flows. To examine the impacts of groundwater recharge on flood inundation and flows in the dry season (October to May), two groundwater recharge scenarios are tested in the Ramganga sub-basin. Increasing groundwater recharge by 35 and 65 % of the current level would increase the baseflow during the dry season by 1.46 billion m³ (34.5 % of the baseline) and 3.01 billion m³ (71.3 % of the baseline), respectively. Analysis of pumping scenarios indicates that 80 000 to 112 000 ha of additional wheat area can be irrigated in the Ramganga sub-basin by additional SSS without reducing the current baseflow volumes. Augmenting SSS reduces the peak flow and flood inundated areas in Ramganga (by up to 13.0 % for the 65 % scenario compared to the baseline), indicating the effectiveness of SSS in reducing areas inundated under floods in the sub-basin. However, this may not be sufficient to effectively control the flood in the downstream areas of the GRB, such as in the state of Bihar (prone to floods),

which receives a total flow of 277 billion m³ from upstream sub-basins.

1 Introduction

Matching water demand with supply in river basins with monsoonal climate is a major challenge. The monsoon-driven seasonal hydrology in India is often associated with floods and droughts, which affects the most vulnerable people of society (women and children, the poor and other disadvantaged social groups), and causes damage to crops and infrastructure. In these basins, upstream storage is generally the preferred solution to buffer the variability of flow and reduce floods downstream (Khan et al., 2014). Traditionally, dams are the major surface water storage structures. However, the construction of large dams requires huge investments, displaces people, submerges forests, and some of the water is lost to non-beneficial evaporation (Pavelic et al., 2012). In contrast, underground aquifers are efficient water reservoirs with minimum evaporative losses and no displacement of people or submergence of land (Bouwer, 2000; Dillon, 2005; Ghayoumian et al., 2007).

For centuries, the utilization of water resources in the Ganges River basin has been severely hampered by substantial seasonal variation in river flows. In the basin, the main source of water is the (southwestern) monsoon rainfall (June to September), and also the snowmelt and ice melt in the Himalaya during the dry season (Sharma and de Con-dappa, 2013). Out of the 1170 billion m³ (billion cubic meters) of water entering the basin, around 500 billion m³ become river flow, while the remainder is returned to the atmosphere through evapotranspiration (World Bank, 2015). The monsoon (between June and September) contributes about

80 % of total annual rainfall and about 80 % of the annual river flow (Revelle and Lakshminarayana, 1975). The rainfall during the rest of the year is low and the river flows, generated mainly through recharged groundwater and snowmelt, are barely sufficient to satisfy the water needs of all the sectors (Huda and Shamsul, 2001). For instance, the estimated average annual flow (1990 to 2008) at the Harding Bridge in Bangladesh (just downstream of the Indian border, with a drainage area of 944 000 km²) was about 340 billion m³ and ranged from 197 to 486 billion m³, whereas flow in the dry season, at the same location, varied from 43 to 63 billion m³.

Extensive flooding in the Ganges River basin, especially in the downstream areas, occurs annually (Mishra, 1997). The major causes of floods in the downstream areas are the shallow groundwater table and high monsoonal rainfall in these areas, and the large runoff generated in the upstream sub-basins. Previous studies (Revelle and Lakshminarayana, 1975; Sadoff et al., 2013) indicated that, due to the limitation of the construction of large surface reservoirs, recharging groundwater beyond the natural level is the best way to control floods downstream. Subsurface storage (SSS) also allows one to meet water requirements during the dry months. Popular belief is that having large dams is the only option to meet the basin’s water storage needs (Onta, 2001). However, contrary to that, the Ganges strategic basin assessment conducted by the World Bank (2012) found that the sustainable use of the basin’s vast groundwater aquifers can store far greater volumes of water compared to the potential of man-made storage in the basin, which is about 130–145 billion m³ (Sadoff et al., 2013). For instance, the mean annual replenishable groundwater in the Ganges basin is about 202.5 billion m³ (Ministry of water resources, 2014). Another study found that the estimated storage available in the shallow alluvial aquifers of eastern Uttar Pradesh and Bihar, which could be utilized in the dry season and naturally recharged in the wet season, is 30–50 billion m³ (SMEC, 2009).

From a purely biophysical perspective, four conditions are necessary to develop sustainable SSS solutions (that involve groundwater recharge beyond the natural levels) to tackle water scarcity and flood damage in the basin:

1. existence of adequate unmet demand (e.g., for agriculture and other uses) to deplete the water pumped from the aquifers in a basin/sub-basin;
2. existence of adequate flows for capture during the monsoon season;
3. existence of extra underground space, which can be created by pumping and depleting groundwater before the onset of the monsoon;
4. ability to actually capture the excess monsoon runoff to recharge that additional space created – naturally (through surface water and groundwater interactions) or artificially (through managed aquifer recharge – MAR).

Amarasinghe et al. (2016) examined the first condition above and estimated unmet demand throughout the basin under two scenarios of irrigation expansion. The main objective of this paper is to examine the second condition above, i.e., assess the potential availability of runoff and the impact of managed groundwater recharge on the river flow. A hydrological model – the Soil and Water Assessment Tool (SWAT) – was used to conduct a hydrological analysis of the sub-basins of the Ganges River basin. This study does not determine whether there is sufficient aquifer storage available to hold the excess runoff, as this requires detailed groundwater aquifer modeling in sub-basins of the GRB. In fact, a comprehensive assessment of the groundwater system in the Ganges is beyond the scope of this work. To the best of the authors’ knowledge, no such work has been done for the whole of the GRB, although this could be done by using the Gravity Recovery and Climate Experiment (GRACE) satellite (Swenson and Wahr, 2006; Morrow et al., 2011; Rodell et al., 2009). Rodell et al. (2009) used GRACE satellite data to estimate the mean rate of groundwater depletion over the Indian states of Rajasthan, Panjab and Haryana as $17.7 \pm 4.5 \text{ km}^3 \text{ y}^{-1}$. Chinnasamy (2017) estimated the groundwater depletion rate over the Ramganga sub-basin located in the northwestern part of the GRB as $1.6 \text{ km}^3 \text{ yr}^{-1}$, and concluded that the depleted aquifer volume can be used to store up to 76 % of the rainfall in the sub-basin. Khan et al. (2014) showed that the subsurface storage created in Uttar Pradesh by pumping groundwater during dry periods can accommodate up to 37 % of the yearly average monsoon flow.

Recharging of runoff to the groundwater aquifer during the monsoon season may have a minimal effect on the downstream flow during the monsoon season. In fact, increased groundwater recharge may increase the contribution of groundwater to the river flow. However, the excess pumping of water from the aquifer can affect the dry season flows. Sadoff et al. (2013) mentioned that using aquifers to store excess water is a national-level alternative for upstream water storage and has the potential to argument dry season flows (although it requires other factors such as appropriate energy pricing and policy environment in conjunction with a well-managed surface water system). Khan et al. (2014) suggest that not withdrawing water from the river during the dry season (which makes up to 50 % of the 28 billion m³ of the annual water withdrawal) in the state of Uttar Pradesh (UP) will increase flow by 25 % in the Ganges at the UP–Bihar boundary. But the authors do not mention how to meet the unmet demand. The reduced surface water pumped can be replaced with increased groundwater pumping (augmented with artificial recharge during the previous wet period). Investigation of the effect of increased groundwater recharge and abstraction on downstream low flows requires conjunctive modeling that couples both groundwater and surface water models. In this study SWAT (which has a simplified groundwater model linked to the surface water model) is used to demonstrate this in the Ramganga sub-basin located in the northwestern part

Table 1. An overview of the main data sets used in this study.

Category	Data	Data source
Topography	Digital elevation model (DEM)	Shuttle Radar Topography Mission (SRTM)
Land use	Land-use map	IWMI database – satellite-based land-use map
Soils	Digital map of soils and soil properties	FAO soil map of the world, 1995
Climate	Rainfall, temperature, relative humidity, sunshine hours, wind speed	Meteorological organization in Bangladesh, re-analysis data, India Meteorological Department (CSFR, 2017)
Hydrology	River discharge	IWMI Water Data Portal

of the GRB. Although this study is a theoretical exercise, it provides a scientific justification for a complete investigation (including field pilot tests) into the plausibility for a well-designed managed aquifer recharge program to enhance the sub-surface storage in the GRB.

2 Methodology

2.1 The model

Many models have been developed (e.g., Eastham et al., 2010; Gosain et al., unpublished data; World Bank, 2012) to study water issues in the Ganges River basin (Johnston and Smakhtin, 2014). However, they are not available to the public. To overcome this restriction and provide the research community with a working hydrological model for the Ganges River basin, the International Water Management Institute (IWMI) has developed a publicly available hydrological model for the basin (Muthuwatta et al., 2014) using the Soil and Water Assessment Tool (SWAT) (Arnold et al., 1998). The model setup files can be downloaded from the website http://waterdata.iwmi.org/pages/model_inventory.php and used in further applications and scenario analyses in a variety of projects.

SWAT is a widely used, semi-distributed conceptual hydrological model developed by the Agricultural Research Service of the United States Department of Agriculture (USDA) over the last 30 years, and is available free of charge as a public domain model (Arnold et al., 1998; Gassman et al., 2007; Sood et al., 2013). The model has been previously being used for number of studies for different watershed scales (e.g., Muttiah and Wurbs, 2002; Ringler et al., 2010; Singh and Gosain, 2011; Sood et al., 2013). The hydrological ability of the model to capture real-world situations is extensively discussed in these articles. Broadly, the SWAT input data can be grouped into five categories: topography or terrain, land use, soil, land use management and climate (Neitsch et al., 2002). SWAT possesses adequate representation of processes governing hydrology and is particu-

larly suitable for application in large river basins. In SWAT, a river basin is subdivided into a number of catchments, so that each catchment has at least one representative stream. Based on unique combinations of soil, land use and slope, the catchments were further divided into hydrological response units (HRUs), which are the fundamental units of calculation. Subdividing the watershed into areas having unique land use, soil and slope combinations enables the model to reflect differences in evapotranspiration and other hydrologic conditions. HRUs allow for a modeling efficiency by lumping pixels with similar land use, soil and slope properties.

SWAT simulates the local water balance of the catchment through four storage volumes – snow, soil profile, shallow aquifer and deep aquifer – based on the soil water balance (Eq. 1):

$$SW_t = SW_0 + \sum_{t=1}^t (R_t - SR_t - ET_t - P_t - G_t), \quad (1)$$

where SW_t is the soil water content minus the wilting-point water content at time t , and R_t , SR_t , ET_t , P_t , and G_t are the daily amounts (in millimeters) of rainfall, runoff, evapotranspiration, percolation and groundwater flow, respectively, at time t . SW_0 is the initial soil water content. The simulated processes include direct runoff (in SWAT direct runoff is termed surface runoff), infiltration, evaporation, transpiration, lateral flow, and percolation to shallow and deep aquifers.

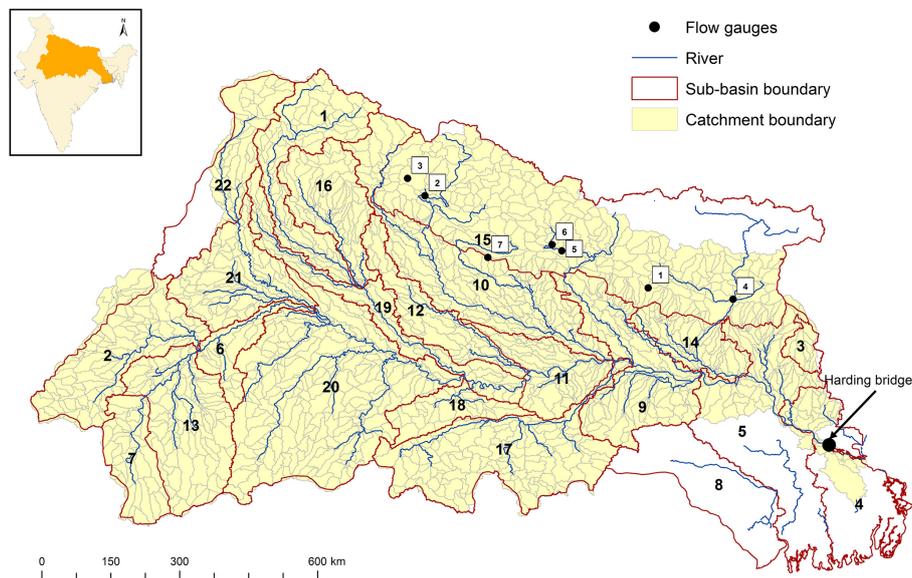
2.2 The data and model setup

The model used in this study was set up using the data sets shown in Table 1. The Ganges River basin was delineated using 3000 ha as the minimum area threshold and has resulted in 1684 catchments (Fig. 1). The area threshold was selected by trial and error in an attempt to represent major tributaries in GRB, while also keeping the SWAT sub-basins to a minimum.

The model was initially developed to study river flow entering Bangladesh. Therefore, the spatial domain of the

Table 2. Model performance indicators for seven locations in the GRB.

Gauge	River	Latitude	Longitude	Period	R^2	NS	RMSE ($\text{m}^3 \text{s}^{-1}$)	Max. flow ($\text{m}^3 \text{s}^{-1}$)
1	Baghmati	27.15	85.49	1981–2006	0.83	0.82	39.7	987.0
2	Karnali	28.96	81.12	1981–2006	0.79	0.61	224.4	2140.7
3	Seti	29.30	80.78	1986–2006	0.76	0.54	92.3	827.4
4	Arun	26.93	87.15	1986–2006	0.63	0.64	446.7	2300.6
5	Kali Gandaki	27.88	83.80	1996–2006	0.75	0.62	280.8	2420.6
6	Kali Gandaki	28.00	83.61	1987–1995	0.58	0.58	261.4	1081.9
7	Kali Gandaki	27.75	82.35	1984–2006	0.76	0.66	293.6	2710.4

**Figure 1.** Sub-basins and catchments of the Ganges River basin (names of the sub-basins are given in Table 3).

SWAT model developed for the Ganges does not entirely cover the areas that belong to West Bengal and Bangladesh. However, this does not affect the current study, as its focus is to assess water availability in the upstream sub-basins of the Ganges River basin.

Figure 1 shows the 22 major sub-basins (Table 3) in the GRB as defined by the Central Water Commission (CWC) of India, which is the main government agency responsible for water resource development and management in the Ganges River basin. Since the focus of this study is to estimate water availability in the sub-basins within India, Nepal is considered one region. The smaller spatial units inside those 22 sub-basins and Nepal are termed “catchments” and were developed using the SWAT interface, as discussed above. The catchments do not completely match with some of the sub-basins due to limitation in SWAT in processing coastal basins.

The model was initially calibrated and validated for the monthly discharge data collated at Harding Bridge. The cal-

ibration period was selected from 1981 to 1990 and the validation period was selected as 1991–2000. The performance indicators, Nash–Sutcliffe efficiency (NS) and coefficient of determination (R^2) were 0.69 and 0.73, respectively, for the calibration period and indicate reasonable agreement between observed and simulated river flow time series. For the validation period, NS and R^2 were 0.75 and 0.81. Additionally, the model simulations were compared with the observed flow data at another seven locations for which the observed data were available. Table 2 presents the model performance indicators for these seven locations. The performance indicators show reasonable values. Further, simulated water balance components seem to be comparable to the results of the other similar studies (e.g., Gosain and Sirinivasan, 2011). For more details on the model setup, including calibration and validation, please refer to Muthuwatta et al. (2014).

Table 3. Runoff of the sub-basins.

Number	Sub-basin	Runoff (billion m ³)			Share of runoff as a percentage of the total	
		Mean	Standard deviation	SR ₇₅	Wet months (June–October)	Dry months (November–May)
1	Above the Ramganga confluence	10.02	5.04	5.48	81.2	18.8
2	Banas	9.89	7.11	3.51	93.8	6.2
3.4	Bangladesh	–	–	–	–	–
5	Bhagirathi and others	–	–	–	–	–
6	Chambal Lower	2.24	1.37	1.23	94.8	5.2
7	Chambal Upper	8.73	3.01	6.60	90.2	9.8
8	Damodar	–	–	–	–	–
9	Gandak and others	16.03	6.57	11.79	86.0	14.0
10	Ghaghara	35.56	17.55	23.34	84.0	16.0
11	Ghaghara confluence to the Gomti confluence	4.72	2.07	3.32	88.3	11.7
12	Gomti	13.64	7.34	9.75	90.8	9.2
13	Kali Sindh and others up to the confluence with Parbati	15.48	6.64	10.51	80.9	19.1
14	Kosi	9.44	3.95	6.81	72.8	27.2
15	Nepal	63.17	11.59	54.44	88.0	12.0
16	Ramganga	15.56	7.79	10.11	82.6	17.4
17	Son	19.50	7.88	14.08	85.1	14.9
18	Tons	6.75	2.47	5.17	88.5	11.5
19	Upstream of the Gomti confluence with Muzaffarnagar	9.38	4.77	5.70	87.8	12.2
20	Yamuna Lower	22.42	10.78	15.21	93.8	6.2
21	Yamuna Middle	4.81	3.70	2.14	78.7	21.3
22	Yamuna Upper	7.19	3.92	4.49	82.7	17.3

2.3 Simulating sub-basin runoff

Annual time series of catchment-scale runoff from 1991 to 2010 were constructed by aggregating daily runoff simulated by SWAT. Next, using geographic information system (GIS) techniques, annual runoff time series were estimated for all sub-basins within the modeled area of the GRB. The study uses the hydrographs of the simulated runoff (SR) to estimate the 75 % dependable runoff (SR₇₅). SR₇₅ is an estimate of the runoff that can be expected in the basin, on average, every 3 out of 4 years, and is considered to be a reliable estimate of water availability for augmenting groundwater storage (Wang et al., 2014).

2.4 Simulating groundwater recharge scenarios in Ramganga

To examine the effect of groundwater recharge on the hydrology such as monthly river flow, the Ramganga sub-basin located in the northwestern part of the basin was selected. The Ramganga sub-basin was selected because it is the first major upstream basin with the typical water resource management challenge of managing seasonal water variability and meeting water demand. The area of the Ramganga sub-basin is about 32 000 km² and it belongs to two administrative dis-

tricts: Uttaranchal and Uttara Pradesh. The important tributaries that flow into the Ramganga River are Kho, Gangan, Aril, Kosi, and Gorra. The surface water potential in the basin is about 18.6 billion m³. The population in the basin is about 20 million. The groundwater recharge was controlled in the SWAT model by changing the curve number (CN). The CN determines the runoff in hydrological models. Reducing the CN in the SWAT increases groundwater recharge.

2.5 Linking river flow to flood inundated areas

The study conducted by Amarnath et al. (2012) developed a data set that used the algorithm based on a number of water and vegetation indices – the Land Surface Water Index (LSWI), the Enhanced Vegetation Index (EVI), the Normalized Difference Vegetation Index (NDVI) and the Normalized Difference Snow Index (NDSI) – on the MODIS 8-day surface reflectance bands to estimate spatial extent and the temporal patterns of flood inundated areas (Amarnath et al., 2012). This data set was used to acquire the maximum flood inundated area for Ramganga. The effect of runoff on maximum flood inundated area in Ramganga was investigated by relating annual values of maximum flood inundated areas to the river flow using logarithmic regression from 2003 to 2010.

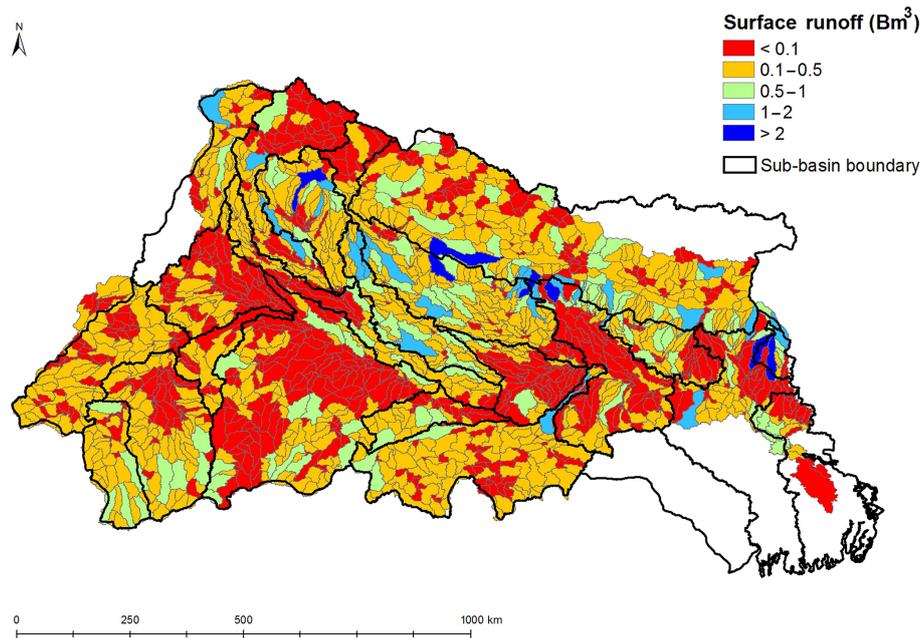


Figure 2. Mean annual runoff of the 1684 sub-basins (1991–2010).

3 Results

3.1 Runoff of the sub-basins

The spatial and temporal distribution of the annual runoff is analyzed to determine the water availability in different sub-basins. River flow includes direct runoff on surface, lateral flow and baseflow from groundwater, which can be captured by diversion or from dams. Direct runoff is calculated in SWAT using the SCS curve number method (Soil conservation service, 1972). In standard hydrological definitions, it is infiltration excess overland flow, which is part of precipitation, that is left after infiltration. It can be captured for MAR before it reaches the stream (in this paper runoff is referred to as the direct runoff calculated by SWAT). Therefore, only the runoff portion was considered for augmenting SSS. Figure 2 shows the simulated catchment-scale mean annual runoff.

The runoff of catchments ranges from less than 0.1 billion m^3 to more than 2.0 billion m^3 . The statistics of the estimated surface runoff for the sub-basins are given in Table 3.

The estimates of mean annual runoff at sub-basin scale range from 2.24 billion m^3 in Chambal Lower (6) to 63.17 billion m^3 in Nepal (15). Additionally, the high standard deviations in Table 4 indicate significant temporal variation within sub-basins. Further analysis shows that runoff in the wet months (June to October) is more than 80 % of the annual runoff in most sub-basins (Table 4, last two columns). This intra- and inter-annual variability of the flows clearly indicates the need for storages to capture the excess runoff during the monsoon season, which could be SSS. For this

analysis, SR_{75} was used to identify the sub-basins that consistently produce higher volumes of runoff. Figure 3 shows the spatial distribution of the SR_{75} of the sub-basins.

The Ghaghara (10) sub-basin and Nepal have, by far, the largest SR_{75} . The Kali Sindh (13), Ramganga (16), Son (17) and Yamuna Lower (20) sub-basins have more than 10 billion m^3 of SR_{75} . The Gandak (9) also produces higher runoff, but the sub-basin is located in the downstream area of the Ganges River basin. Because of the high monsoon runoff, the upstream sub-basins contribute substantially to flooding in the downstream areas of the Ganges River basin.

3.2 Total discharge of the sub-basins

The mean annual discharge from the upstream sub-basins from 2001 to 2010 was estimated and is presented in Fig. 4.

The highest flow of 142.7 billion m^3 to Bihar in the downstream of the GRB comes from upstream of the Gomati confluence to Muzaffarnagar (19), as it gets a large contribution from the Yamuna Lower (20) and Ramganga (16). The second highest flow (78.2 billion m^3) to Bihar comes from the Ghaghara sub-basin (10) and it receives outflows from the western part of Nepal. The mean annual flow to Bihar from the various sub-basins in the Indian part of the Ganges River basin is about 277 ± 121 billion m^3 , and the mean annual rainfall in Bihar is about 123 ± 32 billion m^3 . This indicates that the water volumes received from upstream flows are more than 2-fold the amount of rainfall in Bihar. Flow from the Ghaghara and Yamuna Lower sub-basins is approximately 30 % of the total inflow from the upstream Ganges River basin to Bihar. The contributions from Son, Kali Sindh

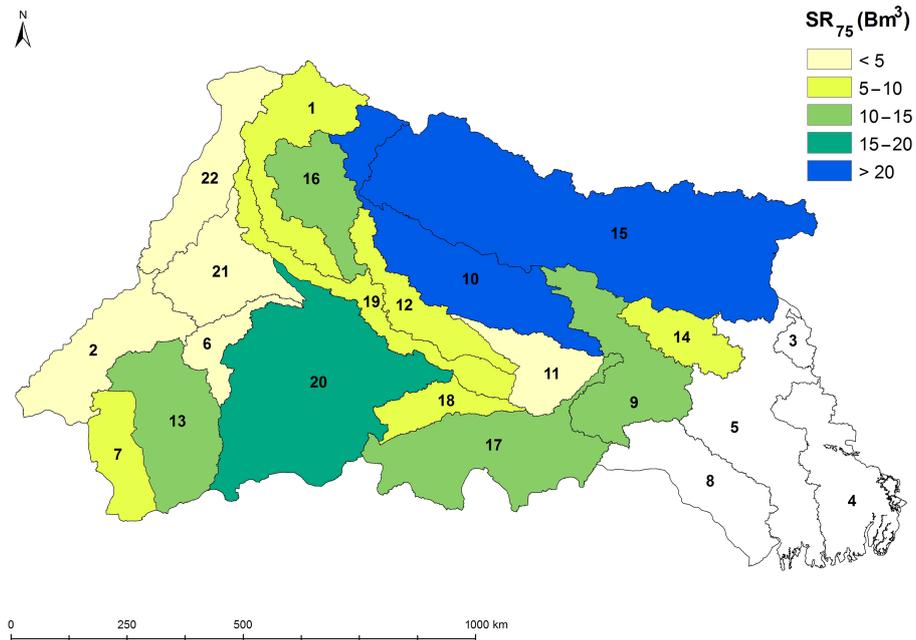


Figure 3. Sub-basin-scale annual dependable runoff (SR₇₅) in the Ganges River basin (1991–2010).

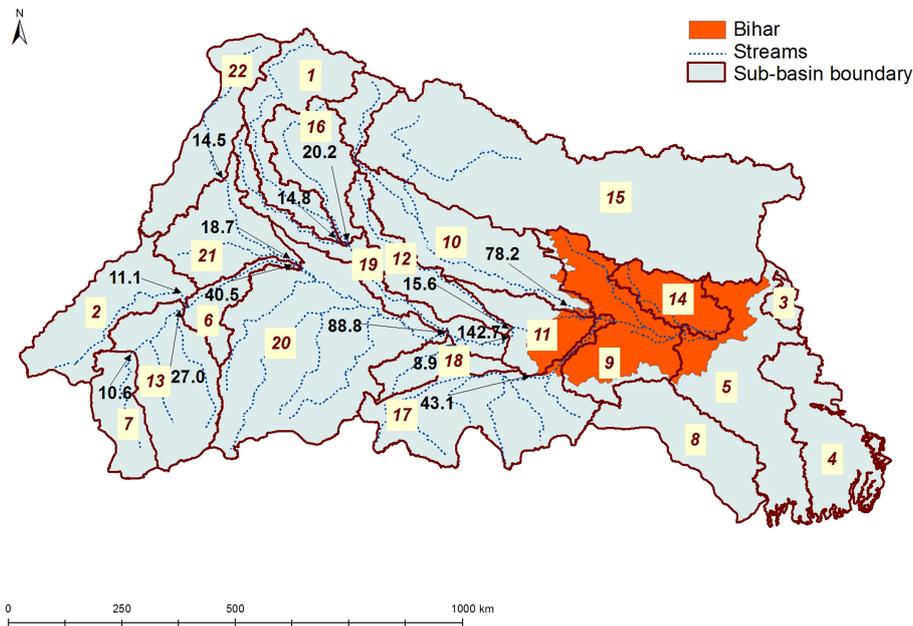


Figure 4. Mean annual outflow (billion m³) from the sub-basins in the Ganges River basin (the numbers in black represent the mean annual outflow, and the numbers in brown in the yellow background represent the numbers of the sub-basins).

and Ramganga are 17, 10 and 7 %, respectively. The estimated discharges at the sub-basin outlets, as shown in Fig. 4, include the contributions from upstream sub-basins and also the contribution of groundwater and runoff to the river flow. Therefore, the values presented in Fig. 4 are significantly higher compared to the surface values presented in Fig. 3.

3.3 Unmet water demand for agriculture

Amarasinghe et al. (2016) estimated the unmet agricultural water demand. Two scenarios were considered in the analysis (Table 4).

- Scenario 1: provide irrigation to the total irrigable area, i.e., increase irrigated area in the Rabi season (Novem-

Table 4. Sub-basin-wise unmet agricultural water demand and the percentage of runoff required to close the unmet demand.

Sub-basin	Unmet demand (billion m ³)		Percentage of the SR ₇₅ required to close the unmet demand	
	Scenario 1	Scenario 2	Scenario 1	Scenario 2
	Above the Ramganga confluence	1.71	2.44	31.2
Banas	1.21	4.09	34.5	116.6
Bangladesh	–	–	–	–
Bhagirathi and others	4.61	15.12	39.1	128.4
Chambal Lower	0.83	1.39	67.7	113.4
Chambal Upper	2.57	5.15	38.9	78.0
Damodar	–	–	–	–
Gandak and others	5.17	7.17	43.9	60.8
Ghaghara	5.11	7.49	21.9	32.1
Ghaghara confluence to the Gomti confluence	3.37	2.89	101.5	87.1
Gomti	2.63	2.83	27.0	29.0
Kali Sindh and others up to the confluence with Parbati	3.9	7.14	37.1	67.9
Kosi	1.03	2.39	15.1	35.1
Nepal	–	–	–	–
Ramganga	2.48	3.28	24.5	32.4
Son	1.92	11.82	13.6	83.9
Tons	0.68	2.34	13.2	45.3
Upstream of the Gomti confluence to Muzaffarnagar	2.93	3.9	51.4	68.5
Yamuna Lower	7.75	18.67	51.0	122.8
Yamuna Middle	3.41	4.72	159.1	220.2
Yamuna Upper	3.72	5.58	82.8	124.2

ber to March) from 26 million ha (current irrigated area in this season) to 30 million ha (irrigable area), and in the hot-weather season (April to June) from 3 million ha (current irrigated area in this season) to 30 million ha (irrigable area), respectively.

- Scenario 2: provide irrigation to the total cropped area. At present, not all cropped area is equipped for irrigation; i.e., irrigable area (30 million ha) is less than the cropped area (35 million ha). Therefore, scenario 2 increases irrigable area and increases irrigated area from 26 to 35 million ha in the Rabi season and from 3 to 35 million ha in the hot-weather (April to June) season, respectively.

As of now, all the sub-basins in the Ganges River basin have substantial unmet water demand for agriculture in the dry season. Therefore, capturing a substantial portion of the runoff during the monsoon months can help close the gap between current supply of water and demand in the dry months, thus increasing agricultural productivity in these sub-basins. Table 4 presents the sub-basin-wise unmet demand and the percentage of dependable runoff required to close the unmet demand.

In the sub-basins, the total unmet demands are 55.03 and 108.4 billion m³ under scenarios 1 and 2, respectively. The values presented in Table 4 show that, for some sub-basins,

annual unmet demand exceeds the annual water availability. In these sub-basins, only a part of the unmet demand can be satisfied by additional underground storage. In some other sub-basins, the unmet demand is less than 30 % of the SR₇₅ of runoff. These sub-basins have the potential to meet all the unmet demand with SSS. In the Ramganga sub-basin, the SR₇₅ of runoff is about 10.1 billion m³, and approximately 83 % of this runoff occurs during the wet season. To meet the maximum unmet agricultural water demand in the Ramganga sub-basin only requires one to capture 33 % of the monsoon runoff.

3.4 Effect of enhanced groundwater recharge and increased pumping on the hydrology

Although runoff is available to store in sub-surface storage (as presented in Tables 3 and 4), it is pertinent to scrutinize the effect of enhanced groundwater recharge and increased pumping on dry season and peak flows in the stream and on downstream water availability. This is demonstrated for the Ramganga sub-basin by simulating hydrology for the baseline scenario and two alternative scenarios: a 35 % increase in groundwater recharge compared to the baseline – S-35; and a 65 % increase in groundwater recharge compared to the baseline – S-65. An increase in groundwater recharge was implemented in the calibrated SWAT model by changing the curve number (CN). Groundwater pumping was im-

Table 5. Mean monthly distribution of river flow and baseflow in the Ramganga sub-basin under different groundwater recharge and abstraction scenarios (BL – baseline scenario; S-35 – 35 % increase in groundwater recharge; S-65 – 65 % increase in groundwater recharge).

Month	Flow			Baseflow (groundwater recharge scenarios)			Additional water requirement C7	Baseflow additional irrigation scenarios S-35			Baseflow additional irrigation scenarios S-65		
	C1	C2	C3	C4	C5	C6		C8	C9	C10	C11	C12	C13
	BL	S-35	S-65	BL	S-35	S-65		100 %	60 %	50 %	100 %	70 %	60 %
Jan	0.24	0.23	0.25	0.16	0.25	0.29	0.27	0.02	0.09	0.11	0.02	0.10	0.13
Feb	0.33	0.28	0.25	0.14	0.23	0.27	0.17	0.07	0.13	0.15	0.10	0.15	0.17
Mar	0.24	0.23	0.24	0.20	0.24	0.29	0.03	0.20	0.22	0.22	0.25	0.26	0.27
Apr	0.12	0.13	0.15	0.17	0.17	0.24		0.17	0.17	0.18	0.20	0.20	0.27
May	0.06	0.07	0.07	0.10	0.11	0.13		0.09	0.10	0.10	0.11	0.11	0.11
Jun	0.87	0.66	0.51	0.05	0.06	0.10		0.06	0.07	0.07	0.08	0.08	0.08
Jul	4.02	3.40	2.90	0.43	0.54	0.80		0.39	0.44	0.45	0.80	0.79	0.79
Aug	6.00	5.51	5.12	1.47	2.07	2.57		2.07	2.07	2.07	2.57	2.57	2.57
Sep	5.33	5.38	5.43	2.24	3.07	3.77		3.07	3.07	3.07	3.77	3.77	3.77
Oct	2.01	2.55	2.99	1.97	2.67	3.25		2.67	2.67	2.67	3.25	3.25	3.25
Nov	0.91	1.23	1.48	1.03	1.39	1.79	0.13	1.26	1.31	1.32	1.66	1.70	1.71
Dec	0.41	0.54	0.64	0.45	0.62	0.97	0.23	0.39	0.48	0.51	0.74	0.81	0.83
Total	20.54	20.20	20.02	8.42	11.41	14.45	0.83	10.46	10.82	10.92	13.55	13.79	13.95

plemented in the SWAT model by removing water from the groundwater storage. The groundwater pumped is assumed to be consumptive use for ET and hence is lost from the system. In Amarasinghe et al. (2016), scenario 2 of unmet agriculture water demand indicated that the agricultural areas in the Ramganga sub-basin could be increased by another 160 000 ha. Thus for this analysis we only consider scenario 2 (from Amarasinghe et al., 2016) of the unmet agriculture water demand. We assume that the additional agriculture area will be wheat, as this crop is predominantly grown during the period of November to March. To estimate water requirements for additional wheat areas from November to March, crop coefficients (k_c) for wheat as obtained from FAO56 (Allen et al., 1998) for similar climatic conditions and crop development stages were used. The Penman–Monteith method served to estimate the daily reference evapotranspiration (ET_0) as required for the crop water requirement estimations. Estimated water requirement for wheat was calculated as 520 mm, which is within the range of recommended water requirements (450–650 mm) for regions with similar settings (see Doorenbos and Kassam, 1979).

Table 5 shows the effect of enhanced groundwater recharge and increased pumping on the baseflow and total streamflow at the main outlet of Ramganga (billion m^3). Columns 1–3 (c1 to c3) present the total streamflow at the main outlet of the Ramganga sub-basin under baseline (BL), S-35 and S-65 scenarios, respectively. Columns 4 to 6 show the simulated monthly baseflow under the three scenarios. Additional water required to expand the irrigated wheat area of 160 000 ha during the period November to March is presented in column 7. The effect of additional pumping under

S-35 and S-65 is presented in columns 8 to 13. Column 8 shows the monthly baseflow if 100 % of the additional area is irrigated by groundwater under the S-35 scenario, while values in columns 9 and 10 are estimated by assuming 60 and 50 % of the 160 000 ha irrigated. Columns 11–13 present the monthly baseflow under S-65 and assume 100, 70 and 60 % of the 160 000 ha irrigated by groundwater, respectively.

Although 85 % of the recharge in Ramganga occurs between July and October, about 80 % of the groundwater contribution (baseflow) occurs during the period August to November (Table 5). The analysis shows reduction of river flow during the high flow months of July, August and, September, for both scenarios as compared to the BL scenario. Under the BL scenario, the streamflow volume at the sub-basin outlet during this 3-month period is 15.34 billion m^3 . It reduces to 14.28 and 13.44 billion m^3 , respectively, when groundwater recharge is increased by 35 and 65 %, respectively (as compared to the baseline scenario). The overall reduction of high flows during this period is 6.89 and 12.37 % for scenarios S-35 and S-65, respectively.

As presented in Table 5, the total increase in baseflow under S-35 and S-65 compared to BL is 3.0 and 6.04 billion m^3 , respectively. The higher baseflow occurs during the 4-month period from August to November. The BL scenario indicates about 6.71 billion m^3 of baseflow during these 4 months, and it increases to 9.19 and 11.37 billion m^3 when groundwater recharge is increased by 35 and 65 %, respectively. During the Rabi season (November to March) the increase in baseflow under S-35 and S-65 is 0.74 and 1.62 billion m^3 , respectively.

Under S-35 scenarios, irrigating 100 % of the additional area would result in reduction of baseflow below the BL scenario during December to February. However, as presented in Table 5, for scenario S-35, additional irrigation to cover 50 % of the new wheat area would still maintain the baseflow above the BL level other than in January. Irrigating 60 % of additional irrigated wheat areas would reduce the baseflow volumes below the BL levels in January and February. Results, further, indicate that under the S-65 scenario it will be possible to supply irrigation to 70 % of the additional irrigated area without reducing the volumes of baseflow simulated in the BL scenario. When it is 70 % of the additional irrigated area, baseflow will reduce by a negligible amount in January.

When water balance is considered, the summation of total baseflow under abstraction scenarios and additional water requirement must be equal to the total baseflow under a no pumping scenario. For instance, the sum of the totals of C8 and C7 must be equal to the total of the C5 column. However, we found some differences which are negligible. In this case, the error is about 0.12, which is about 1.05 % compared to the total baseflow presented in C5. The water balance errors under the remaining five abstraction scenarios also range from 0.0 to 0.8 %. We presume that these small differences are due to changes in other hydrological process such as changes in soil moisture and evapotranspiration as a result of increased groundwater infiltration.

3.5 Effect on floods

The relationship between the simulated maximum monthly river flow and the maximum flood inundated areas in Ramganga is shown in Fig. 5. The horizontal axis represents simulated maximum monthly river flow during each year from 2003 to 2010 at the Ramganga outlet. The vertical axis shows the maximum flood inundation areas estimated based on the satellite images in the corresponding year (Amarnath et al., 2012).

The coefficient of the determinant (R^2) indicates a strong correlation between the area under floods and the annual runoff, and this implies that the maximum monthly runoff explains more than 70 % of the variation in maximum flood inundated area. The mathematical relationship between maximum flood inundated area and the runoff is given in Eq. (2):

$$\text{Maximum flood inundated area} = 568.7 \times \ln(\text{Flow}) - 356.2. \quad (2)$$

The maximum monthly flow in Ramganga of about 6.0 billion m^3 in August (Table 5) has a corresponding flood inundated area of about 660 km^2 . Reduction of peak flow to 5.5 billion m^3 (35 % groundwater recharge scenario) would reduce the flood inundated area by about 6.9 %. Similarly, the reduction of flood inundated areas compared to the baseline scenario is about 13.0 % for the 65 % groundwater recharge scenario. For this scenario, the reduced outflow (in August) from the basin is about 15 %. This analysis shows the poten-

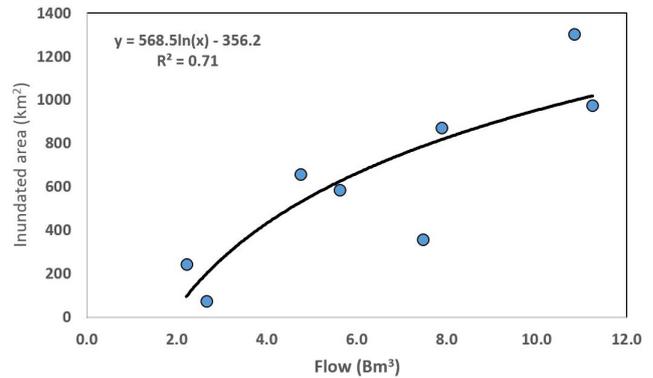


Figure 5. Relationship between annual maximum flood inundated area and the maximum monthly river flow in Ramganga.

tial impacts of enhanced sub-surface storage on the flooding in the Ramganga sub-basin located in the upstream. The volume of inflow in the Ramganga is negligible compared to the inflow received by the areas such as Bihar in the downstream. Therefore, to understand the potential impacts of SSS on flooding in the GRB, further research is required to investigate the effect of SSS on control of floods in the downstream areas.

4 Discussion

Water availability and demand analysis conducted in the Ganges River basin show that there is a substantial mismatch between water demand and supply. For instance, estimated unmet annual water demand for agriculture in the GRB (based on the two scenarios discussed above) ranges from 55.03 to 108.4 billion m^3 , while annual total runoff generated in the basin is about 298 ± 99 billion m^3 , of which 80 % occurs during the monsoon months. In this situation, strategies must be formulated to manage available water in the GRB in a more productive manner. One management option discussed in this paper is using SSS. Augmenting SSS is important in securing downstream water availability for ecosystems and other uses such as agricultural, domestic and industrial.

A thorough analysis of water resource management options requires knowledge of the spatial and temporal distribution of water availability and a substantial amount of hydrological data. In most cases, such data are not publicly accessible. Thus remote sensing and models are helpful in filling in gaps where data are not available. Models are also helpful in analyzing the impact of SSS without making large financial investments. As presented in the results section, SWAT model was calibrated using only flow data and the model performance indicates acceptable results. However, the model calibrated for multiple water balance components would have provided more trustworthy simulations. Other observed data such as actual evapotranspiration and

soil moisture could have made the model more robust, but such data do not exist (although satellite products are there).

Results of the SWAT model demonstrate its capability in estimating the spatial and temporal water availability in the sub-basins of the GRB and in assessing the effect of augmenting SSS on the hydrology of the basin. In all the scenarios, augmenting SSS does not show much difference in total annual flow from Ramganga, but there are intra-year changes. There is reduction in flow during the peak season but an increase in flow during the dry season. This indicates that augmenting SSS can help in flood reduction while improving water availability for crops in the dry season. For the excess irrigation scenarios considered, 80 000 to 112 000 ha of additional agriculture land can be irrigated by groundwater without affecting the baseflow in the basin. There still remain some limitations in this study, mainly due to the limitations with the model such as unavailability of the model in handling of groundwater depth and no detailed linkages between surface water and groundwater (since SWAT is predominantly a surface water model).

Flood inundated areas based on satellite remote sensing data (provided by another study) allowed us to investigate the impact of SSS on downstream floods. However, the relationship established between flood inundated areas and the river flow was only for Ramganga, and further investigations are required to understand how SSS will impact large floods in the downstream part of the basin.

5 Conclusions

Creating additional SSS beyond the current levels in the Ganges River basin can simultaneously enhance water supply and control downstream floods. The sub-basin-wise mean annual runoff ranges from 2.24 to 35.56 billion m³, and the contribution of runoff from Nepal is about 63 billion m³. Several sub-basins in the Ganges River basin produce sufficiently high dependable annual runoff that can be stored underground and used during the dry season. For instance, annual runoff in each of the five sub-basins in the upstream of the Ganges River basin is more than 10 billion m³, which is about 30 % of total runoff generated in the GRB. Comparison of sub-basin-wise runoff with the estimated unmet water demand indicates that capturing only a portion of the wet-season runoff would be sufficient to provide water to irrigate all the irrigable land in the dry months. Sub-basin-wise river flow analysis in the GRB shows that approximately 30 % of the upstream flow to Bihar comes through the Ghaghara and Yamuna Lower sub-basins. This runoff contributes to the recurrent floods in Bihar.

A case study based on Ramganga indicates that increasing 35 and 65 % groundwater recharge compared to the baseline scenario may reduce the peak monthly flow by about 6.8 and 12.3 %, respectively. Further, the results indicate that the dry season flow (October to May) can increase by 21 and 40 % in

these two scenarios before meeting unmet demand by pumping.

More than 70 % of the variations of flood inundated areas in the Ramganga sub-basin can be explained by the maximum monthly river flow values. By increasing groundwater recharge by 35 and 65 % during the peak flow months the flood inundated area can be reduced by about 6.6 and 8 %, respectively.

This study focused on spatio-temporal water availability and the impacts of SSS on the hydrology in the GRB. Pumping scenarios simulated by the SWAT model indicated that additional wheat areas in the Rabi season could be irrigated by the increased SSS under a 35 % increase in groundwater recharge and a 65 % increase in groundwater recharge scenarios.

This study only discusses the surface water availability for SSS, without going into details regarding suitability of recharge areas. A detailed analysis of the soil, topographic and geological characteristics is required to determine the suitable areas for groundwater recharge.

Data availability. Sources of the data used in this study are shown in Table 1. The SWAT model setup files can be found from the website: http://waterdata.iwmi.org/pages/model_inventory.php.

Part of the daily meteorological data sets used here were obtained from Indian Meteorological Department (IMD) and are not available to the public as far as the authors know. Therefore, the authors cannot provide them. The other meteorological data were downloaded from The National Centers for Environmental Prediction (NCEP) HYPERLINK <http://rda.ucar.edu/pub/cfsr.html> Climate Forecast System Reanalysis (CFSR). This data set is freely available <https://globalweather.tamu.edu/> (CSFR, 2017).

Competing interests. The authors declare that they have no conflict of interest.

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