

## Comparison of the results by the present method and the fourth order Runge-Kutta method

For comparing the result obtained by the present method and that by the fourth order Runge-Kutta method, we take the Example 5.8 in the book of K. Subramanya (2009). The book title is “Flow In Open Channels.” The followings show: (1) the content of Example 5.8, (2) the numerical code **for Matlab** by using the present method (Jan & Chen’s method), (3) the numerical code for Matlab by using the standard fourth order Runge-Kutta method, and the comparison of the results by the present method and fourth order Runge-Kutta method.

### Example 5.8

A river 100 *m* wide and 3.0 *m* deep has an average bed slope of 0.0005. Estimate the length of GVF profile produced by a low dam which raises the water surface just upstream of it by 1.50 *m*. Assume the channel cross-section is rectangular, the Manning coefficient *n* equals 0.035, and the normal depth of flow is 3.0 *m*. Find the GVF profile (M1 curve).

### Numerical code for Matlab by the present method (Jan & Chen’s method) :

```
clear all;clc;tic
%-----Basic Information-----
h0=3;           % Normal depth
S0=0.0005;     % Average bed slope
n=0.035;       % Manning coefficient
h1=4.5;        % Water depth (m) at boundary condition, x=0
g=9.81;        % Gravitational acceleration (m/s2)
q=h0(5/3)*S0(1/2)/n; % Unit-width discharge for wide channels by Manning equation (m3/s/m)
hc=(q2/g)(1/3);    % Critical depth (m)
N=10/3;        % Hydraulic exponent for uniform-flow computation
M=3;           % Hydraulic exponent for critical-flow computation in wide channels
hEND=3.03;     % Upstream water depth (m)
xs=0;          % Boundary condition at which the start of GVF calculation
u1=h1/h0;      % Dimensionless water depth at downstream BC x=0
XS=xs*S0/h0;  % Dimensionless longitudinal coordinate
ld=hc/h0;      % Dimensionless critical depth
a=1;           % Parameter used in Gaussian hypergeometric function
b1=-1/N;       % Parameter used in Gaussian hypergeometric function
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c1=b1+1; % Parameter used in Gaussian hypergeometric function
b2=(M-1)/N; % Parameter used in Gaussian hypergeometric function
c2=b2+1; % Parameter used in Gaussian hypergeometric function
z1=u1^(-N); % Dimensionless water-depth parameter used in GHF
hypergeomBF1= hypergeom([a,b1], c1, z1) % GHF-1 at BC
hypergeomBF2= hypergeom([a,b2], c2, z1) % GHF-2 at BC

%-----M1 Curve calculated by Jan &Chen's method -----
for i=1:100000 % Upper-limit of calculation steps
    h(i)=h1-0.01*(i-1); % Water depth (m) at a specified longitudinal location to be determined
    u(i)=h(i)/h0; % Dimensionless variable of water depth
    z(i)=u(i)^(-N); % Dimensionless variable used in GHF
    if y(i)<hEND % Condition for the end of GVF calculation
        break
    end
end
Xx0(i)=XS+(u(i)*hypergeom([a,b1], c1, z(i))-u1*hypergeomBF1)+
ld^M*(u(i)^(-M+1)*hypergeom([a,b2], c2, z(i))-u1^(-M+1)*hypergeomBF2)/(M-1);
    Xx(i)=Xx0(i)*h0/(S0); % transfer the longitudinal coordinate to be dimensional
    if h(i)==hEND % Condition for the end of GVF calculation
        break
    end
end
end
h=h(1:length(Xx)); % Let storage length for water depth equal that for distance Xx
toc

```

### **Numerical code for Matlab by the standard fourth order Runge-Kutta method :**

```

clear all;clc;format long;tic
%-----Basic Information-----
h0=3; % Normal depth (m)
S0=0.0005; % Average bed slope
n=0.035; % Manning coefficient
h1=4.5; % Water depth (m) boundary condition : x=0
g=9.81; % Gravitational acceleration (m/s2)
q=h0^(5/3)*S0^(1/2)/n; % Unit-width discharge for wide channels by Manning equation (m3/s/m)
hc=(q^2/g)^(1/3); % Critical depth (m)
B=100; % Width of a rectangular channel (m)

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Q=q*B; % Flow discharge in a rectangular channel (m3/s)
T=B; % Top width of a rectangular channel (m)
dx=-1; % Distance interval (m)
L=-8644; % Computation end of longitudinal distance (m)
nn=L/dx; % Number of calculation steps
h=[]; % Water depth (m) for the longitudinal datum
%-----Calculation by Standard Fourth Order Runge-Kutta method solved by Matlab -----
for i=1:nn
    LL(i)=dx*i; % Storage for longitudinal distance length
%-----K1-----
    A1=B*h1; % Cross-section area of flow in a rectangular channel (m2)
    P1=2*h1+B; % Wetted perimeter of flow in a rectangular channel (m)
    R1=A1/P1; % Hydraulic radius (m)
    V1=Q/A1; % Average velocity (m/s)
    Sf1=(V12)*(n2)/(R1(4/3)); % Energy slope evaluated by Manning equation
    F1=(S0-Sf1)/(1-Q2*T/(g*A13)); % Function in the basic differential equation of GVF
    K1=dx*F1; % First parameter (m)
%-----K2-----
    A2=B*(h1+K1/2); % Cross-section area of flow in a rectangular channel (m2)
    P2=2*(h1+K1/2)+B; % Wetted perimeter of flow in a rectangular channel (m)
    R2=A2/P2; % Hydraulic radius (m)
    V2=Q/A2; % Average velocity (m/s)
    Sf2=(V22)*(n2)/(R2(4/3)); % Energy slope evaluated by Manning equation
    F2=(S0-Sf2)/(1-Q2*T/(g*A23)); % Function in the basic differential equation of GVF
    K2=dx*F2; % Second parameter (m)
%-----K3-----
    A3=B*(h1+K2/2); % Cross-section area of flow in a rectangular channel (m2)
    P3=2*(h1+K2/2)+B; % Wetted perimeter of flow in a rectangular channel (m)
    R3=A3/P3; % Hydraulic radius (m)
    V3=Q/A3; % Average velocity (m/s)
    Sf3=(V32)*(n2)/(R3(4/3)); % Energy slope evaluated by Manning equation
    F3=(S0-Sf3)/(1-Q2*T/(g*A33)); % Function in the basic differential equation of GVF
    K3=dx*F3; % Third parameter (m)
%-----K4-----
    A4=B*(h1+K3/2); % Cross-section area of flow in a rectangular channel (m2)
    P4=2*(h1+K3/2)+B; % Wetted perimeter of flow in a rectangular channel (m)
    R4=A4/P4; % Hydraulic radius (m)
    V4=Q/A4; % Average velocity (m/s)

```

```

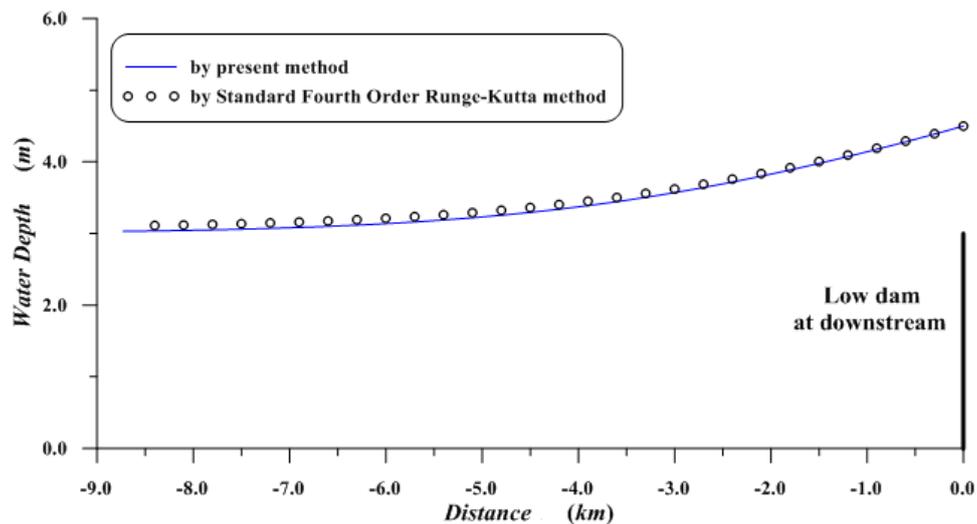
Sf4=(V4^2)*(n^2)/(R4^(4/3));           % Energy slope evaluated by Manning equation
F4=(S0-Sf4)/(1-Q^2*T/(g*A4^3));        % Function in the basic differential equation of GVF
K4=dx*F4;                               % Fourth parameter (m)

h2=h1+(K1+2*K2+2*K3+K4)/6;             % Water depth (m) at the subsequent step
h1=h2;                                   % Renew water depth (m) for next-step calculation
h=[h h2];                                % Storage length for water depth

end
toc

```

### Comparison of the results by the present method and fourth order Runge-Kutta method



The numerical error in water depth obtained by the standard fourth order Runge-Kutta method is about 2% at the longitudinal distance  $x = -8$  km.